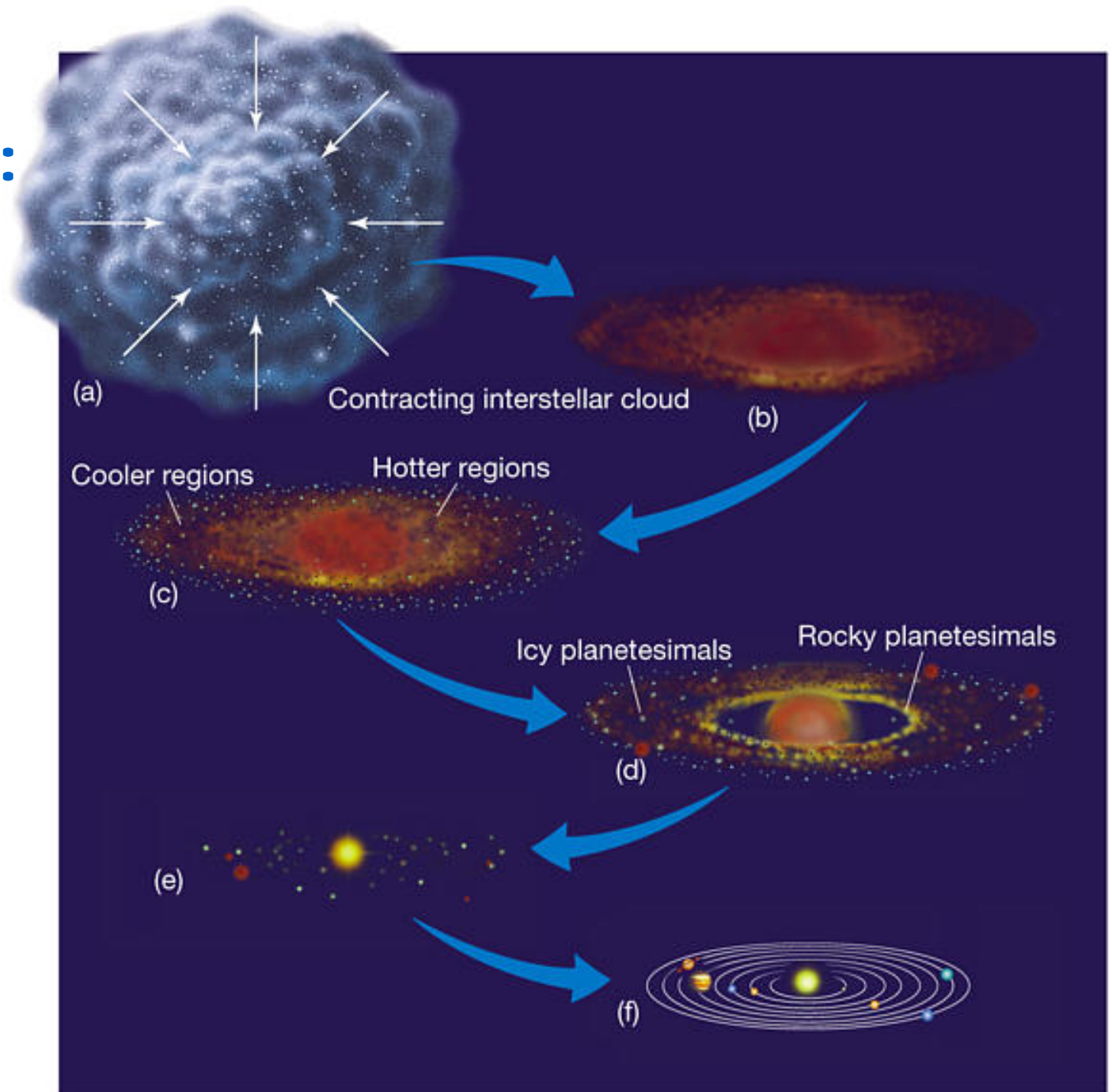


The background of the slide is a detailed digital illustration of a protoplanetary disk. A bright, young star is at the center, surrounded by a glowing disk of gas and dust. In the foreground, a ringed planet, similar to Saturn, is shown in profile. The scene is set against a dark, star-filled space. The text is overlaid on the upper left and lower right portions of the image.

*Theories of planet
formation and migration*

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Planet Formation: Overview

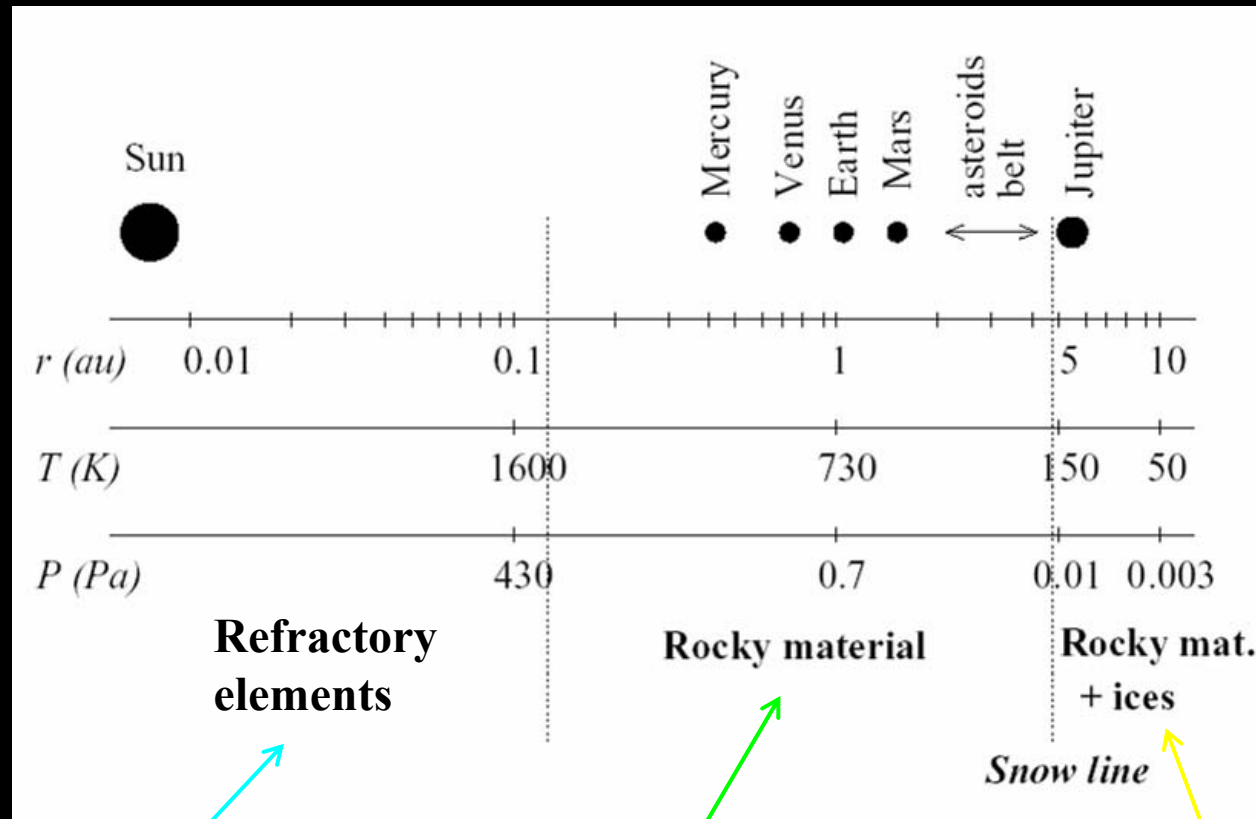


α -Disk Models

$$dM/dt = 10^{-7} M_{\odot} / \text{yr}$$

$$\alpha = 10^{-2}$$

(Papaloizou & Terquem 1999)

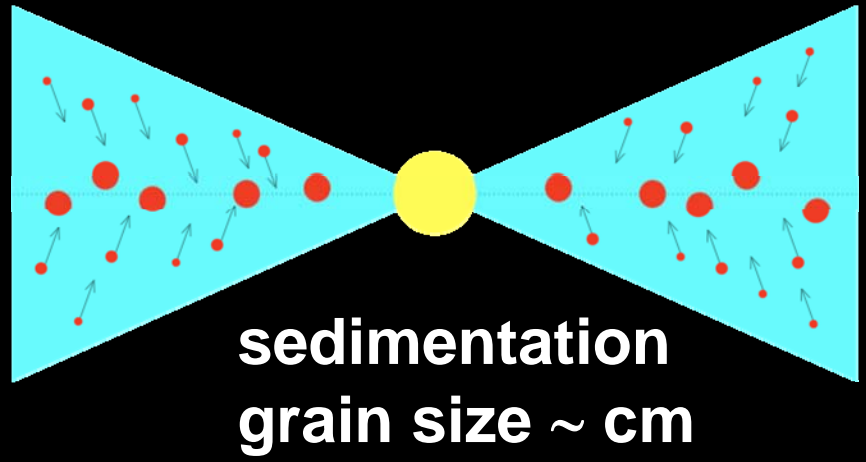


Ca and Al
Oxydes

Solids containing Fe, Mg,
Si and S combined with
themselves and with O

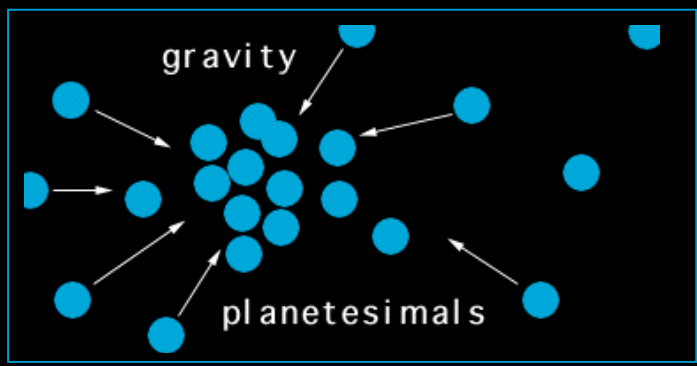
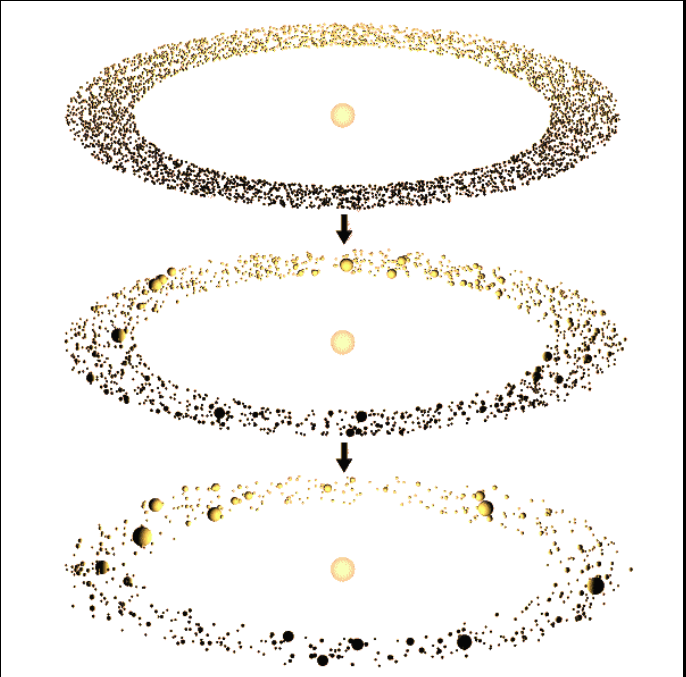
H_2O , CH_4 , NH_3 ,
 CO_2 , CO
Gaseous Hydrates

Formation of terrestrial planets/cores



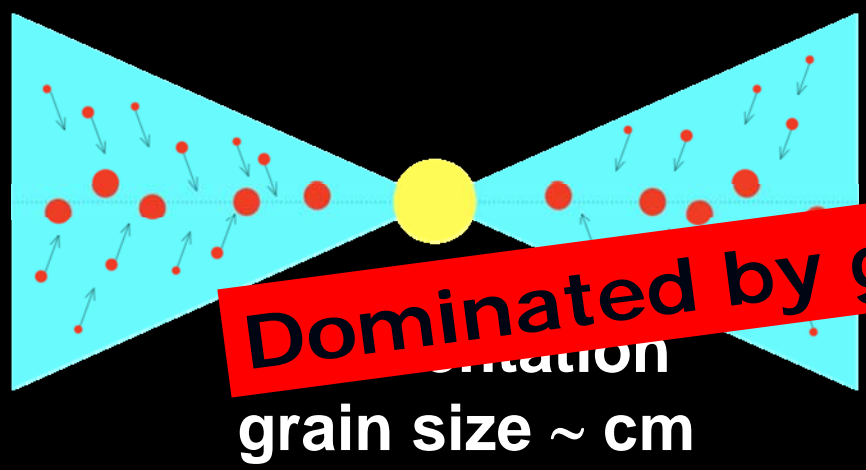
cm → 0.1 - 1 km

(Tanaka et al. 2000)



1 km → few 100 km,
i.e. $10^{-3} - 10^{-2} M_{\oplus}$

Formation of terrestrial planets/cores



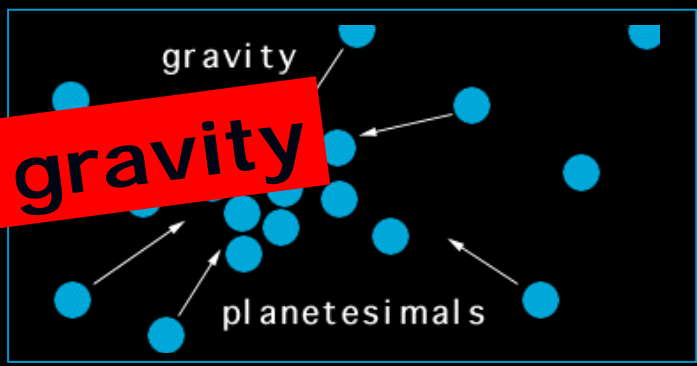
Dominated by grain-gas interaction



cm → 0.1 - 1 km

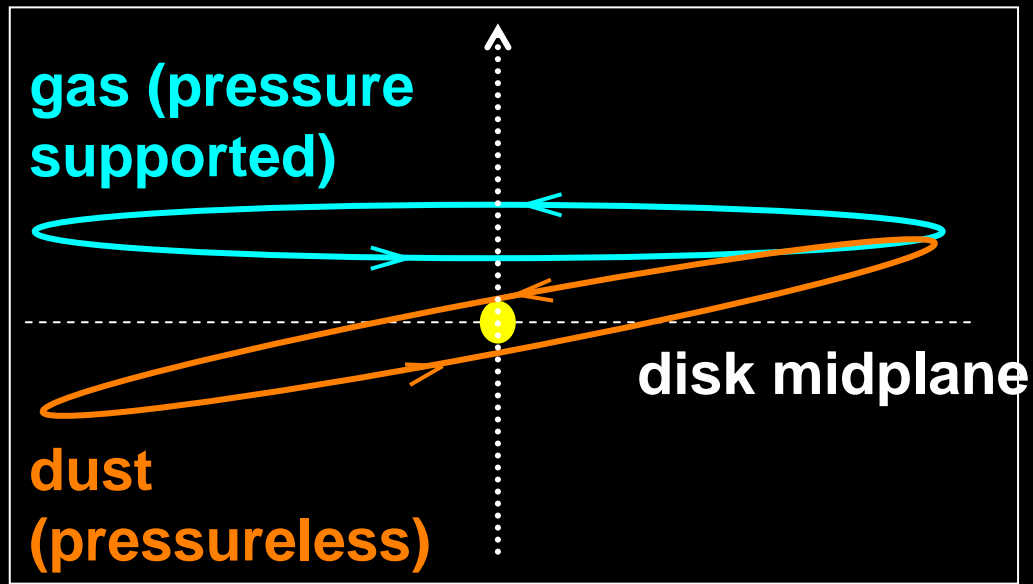


Dominated by gravity



1 km → few 100 km,
i.e. $10^{-3} - 10^{-2} M_{\oplus}$

Sedimentation



In the rotating frame, the dust **oscillates around the disk midplane** with the frequency = orbital frequency

Drag between the dust and the gas → **damping** of the oscillations ($\tau_{\text{damp}} \sim 10 \text{ s}$)

→ **Sedimentation toward the disk midplane**

Collisions between grains → **growth to ~ cm-m size**

Sedimentation timescale ~ **10^5 years at 1 au** with no turbulence

Radial drift

Dust: not pressure supported → **Keplerian** velocity

Gas: pressure supported → **sub-Keplerian** velocity

→ Drag → the dust drifts inward

(Weidenschilling 1977)

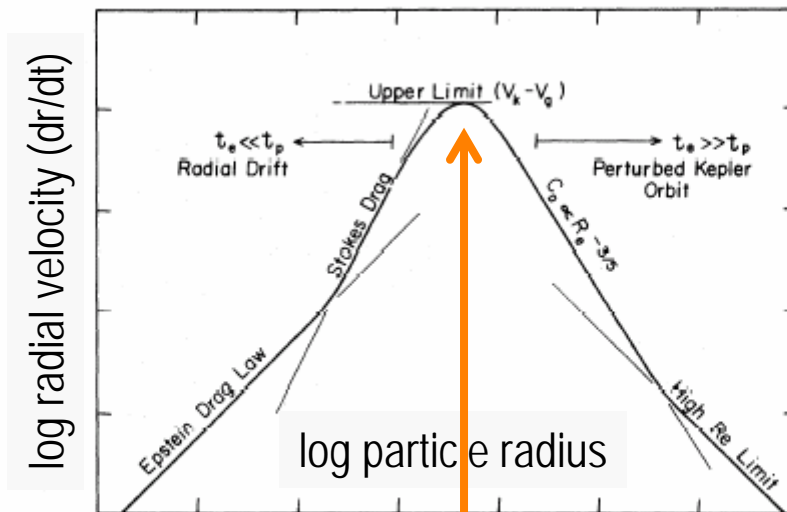


Figure 1. Radial velocity versus particle size (schematic). The shape of the curve is determined by the drag laws, but the peak value depends only on the nebular structure.

meter

Formation of planetesimals

Growth from cm-m to 0.1-1 km:

???

but has to be **fast** since **radial drift** most efficient ($\tau \sim 100$ years, or 10^3 years for collective drag) for m-sized bodies.

Solution = Turbulence?

Particles concentrate at pressure maxima

Streaming instabilities \rightarrow reinforce the concentration

Growth faster than the drift (Johansen et al.)

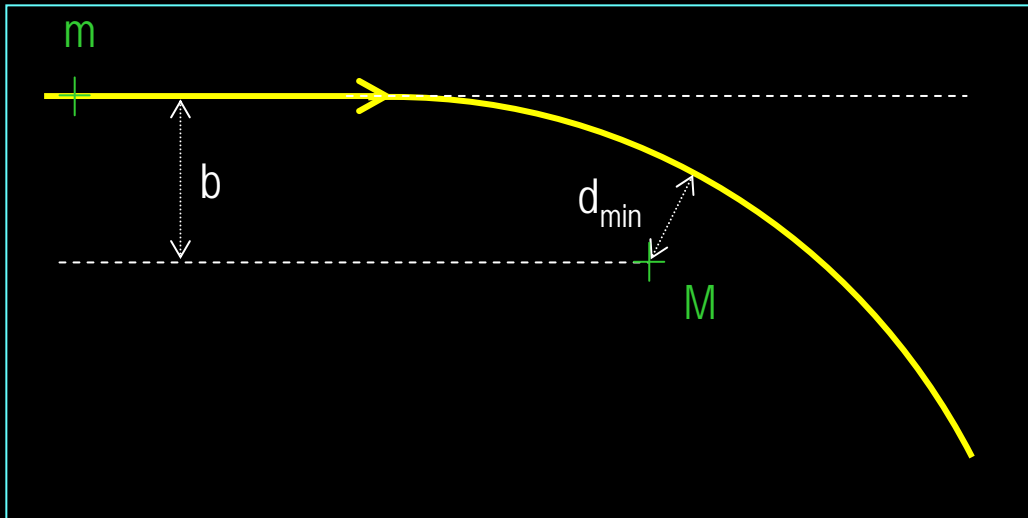
Protoplanet formation

0.1-1 km \rightarrow 100 km (10^{-2} - $10^{-3} M_{\oplus}$): « particles in a box »

collisions and sticking \rightarrow growth

runaway accretion

Runaway accretion



Energy: $E^* = \frac{1}{2}\mu v^2 - GMm/r$

Angular momentum:

$$\mathbf{L}^* = \mathbf{r} \times \mu \mathbf{v}$$

$$L^* = cst \rightarrow \mu b v_{-\infty} = \mu d_{\min} v(d_{\min})$$

$$E^* = cst \rightarrow \frac{1}{2}\mu b v_{-\infty}^2 = \frac{1}{2}\mu v^2(d_{\min}) - GMm/d_{\min}$$

$$\Rightarrow \sigma_{col} = \pi b^2 = \underbrace{\pi (r_m + r_M)^2}_{\sigma_{geom}} \left[1 + \frac{M+m}{M} \frac{v_e^2}{v_{-\infty}^2} \right]$$

v_e : escape velocity at the contact point

σ_{geom}

“gravitational focussing” $\sim 10^3$!

Si $r_M \gg r_m$, $\sigma_{col} \propto r_M^2 \rightarrow$ “runaway accretion”

Runaway growth

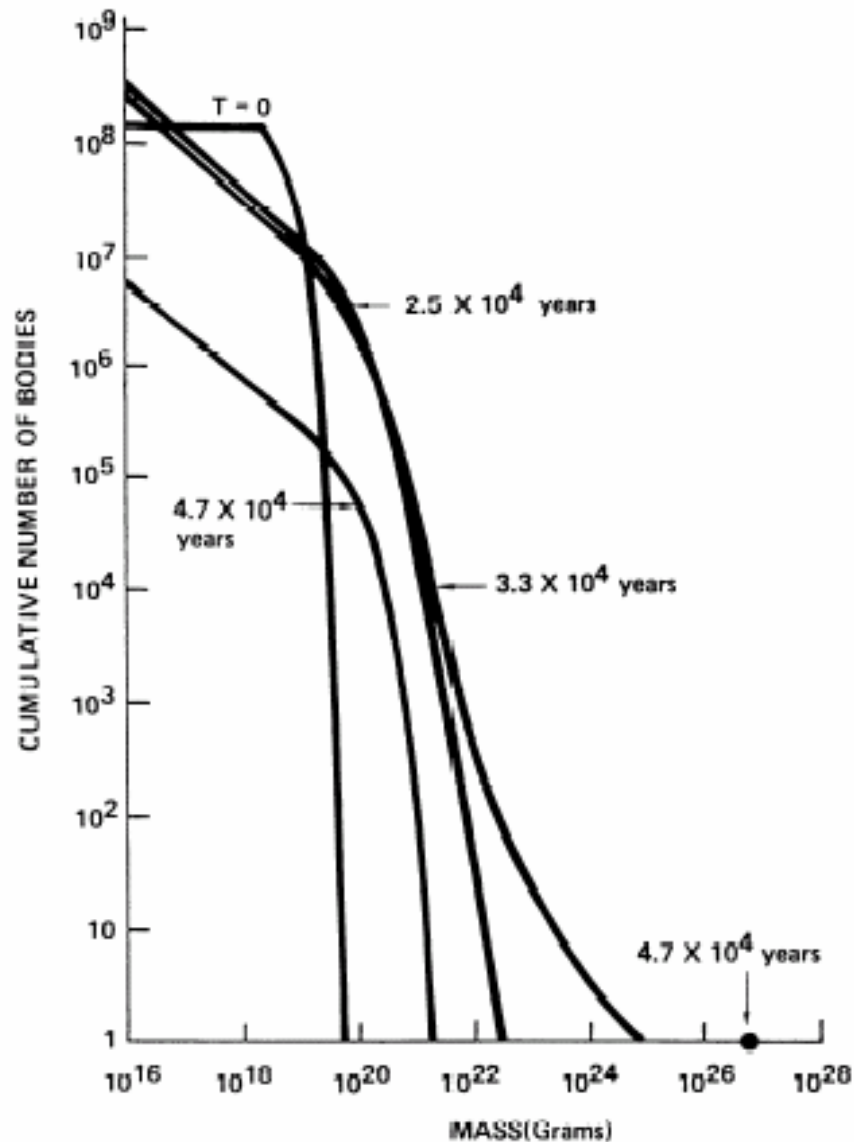
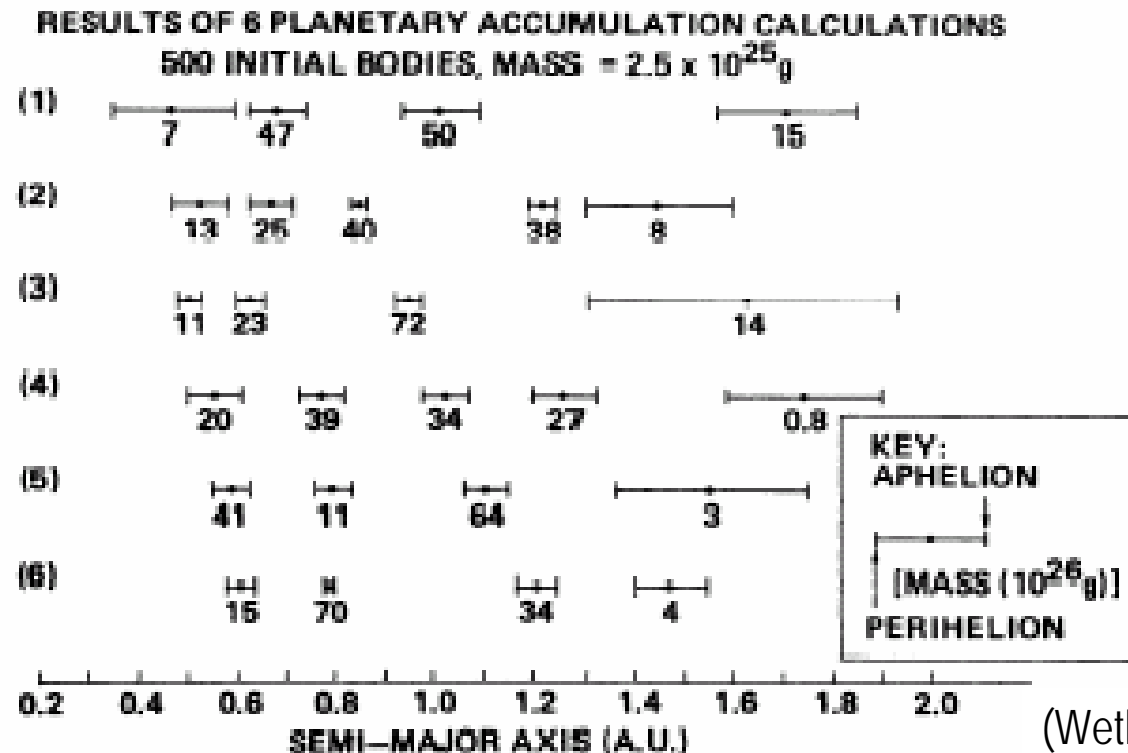


Figure 3 The evolution of the size distribution of a swarm of planetesimals distributed between 0.99 and 1.01 AU using the velocity evolution equations of Stewart & Wetherill (1988). This simulation includes fragmentation, a reduction of gravitational perturbations of runaway bodies from the uncorrelated encounters approximation, and the 3-body gravitational enhancement in accretion cross-sections for low velocity bodies. Note the rapid runaway growth of the largest bodies, with the most massive planetary embryo becoming detached from the remainder of the swarm. Figure adapted from Wetherill & Stewart (1989), courtesy G. Wetherill.

(Wetherill & Stewart 1989
Lissauer 1993)

Planet / core formation

N-body simulations

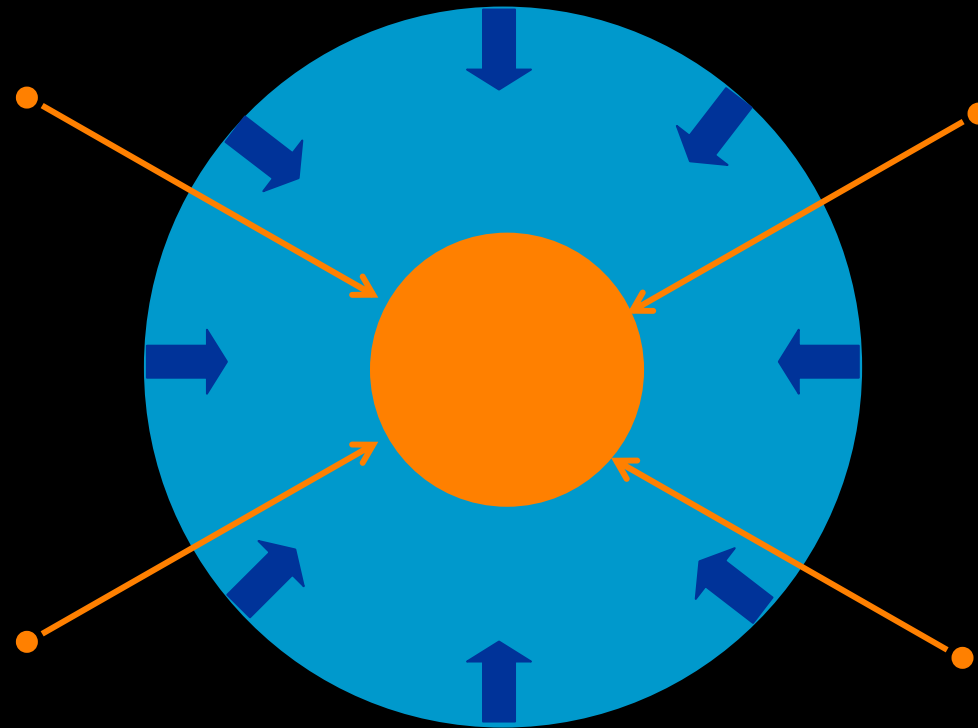


(Wetherill 1989)

Figure 4 Final outcome of 6 terrestrial planet accumulation calculations by Wetherill (1988). The simulations began with 500 bodies each of mass 2×10^{25} g. The semimajor axes of the final planets are indicated by points; the line through each point extends from the perihelion to the aphelion of the planet. The numbers under each point indicate the final mass of the body in units of 10^{26} g.

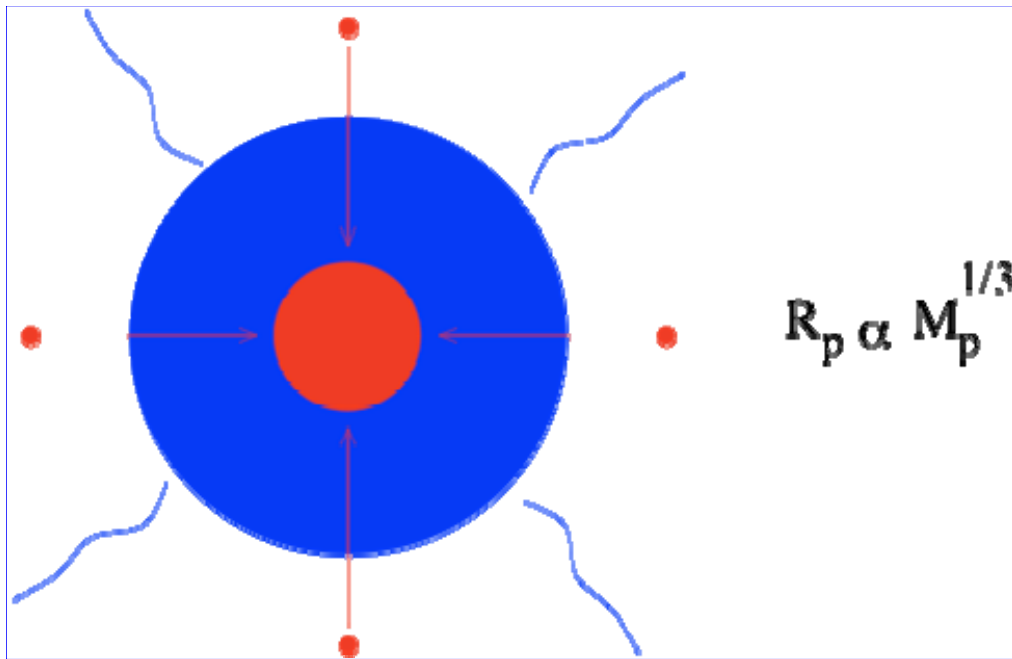
Giant planet formation

Capture of a gaseous envelope:



→ Critical core mass

Capture of an envelope



Energy sources:

(S1) accretion of planetesimals

(S2) gravitational contraction of the gas

Energy losses:

(P1) Radiative transport

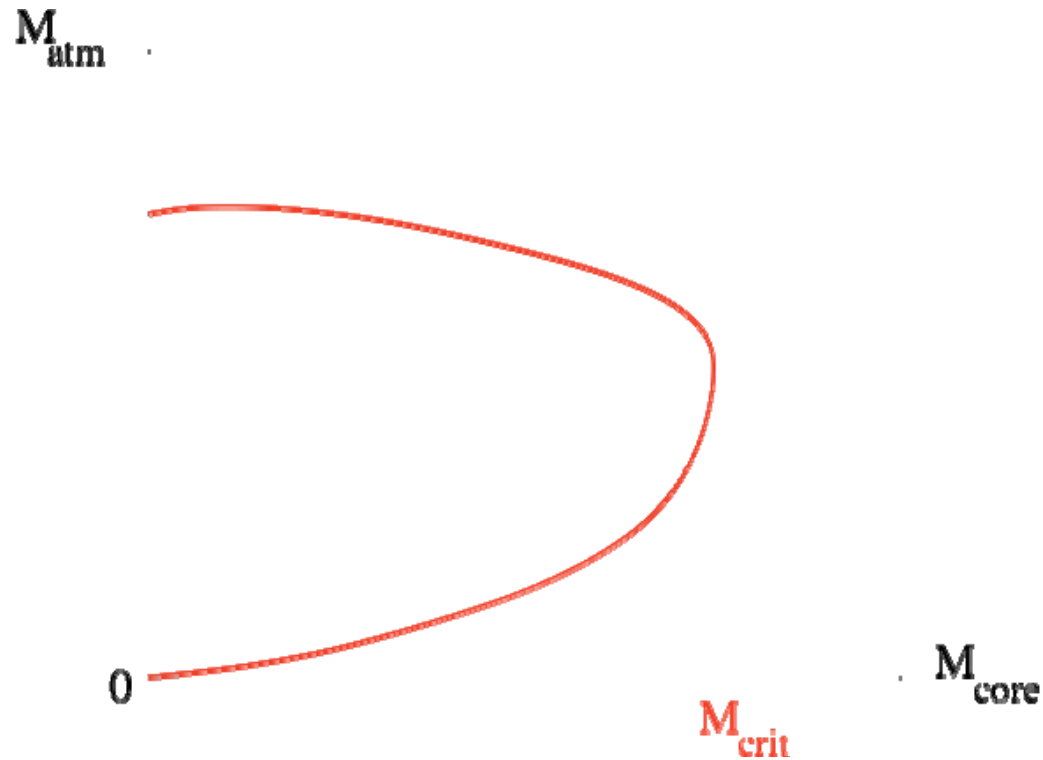
(P2) Convective transport

Energy conservation: $(S1) + (S2) = (P1) + (P2)$

If $(S1) \gg (S2)$: atmosphere at thermal and hydrostatic equilibrium

(Perri & Cameron 1974, Mizuno 1980)

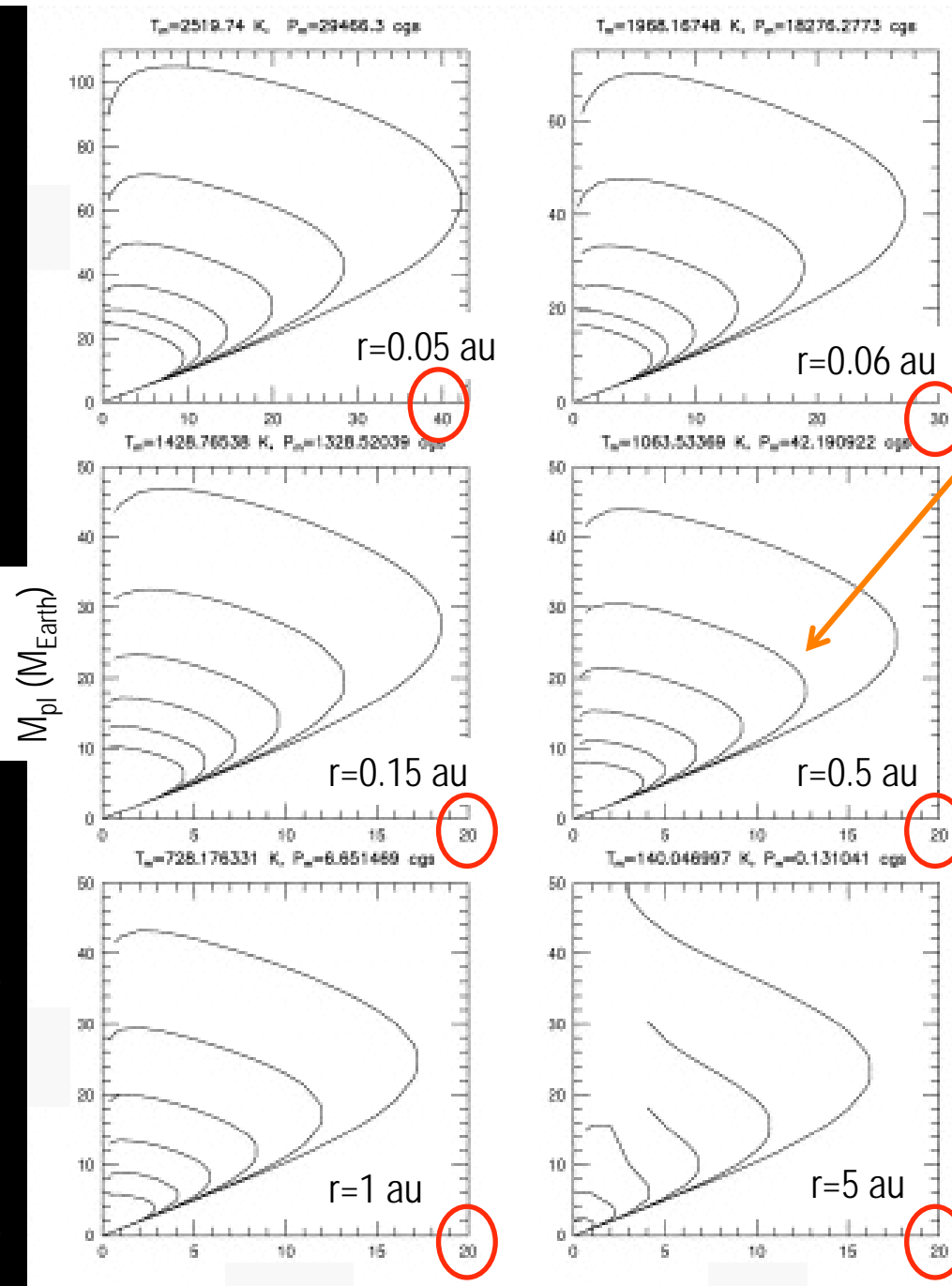
Capture of an envelope



- $M_{\text{core}} < M_{\text{crit}}$: (S1) large enough to support the atmosphere
- $M_{\text{core}} > M_{\text{crit}}$: (S1) not large enough
→ collapse of the atmosphere

M_{crit} increases with dM_{plan}/dt

Giant planet formation



M_{pi} (M_{Earth})

M_{core} (M_{Earth})

$$dM_{plan}/dt = 10^{-7} M_{\oplus}/yr$$

$$M_{crit} = 10 - 15 M_{\oplus}$$

BUT...

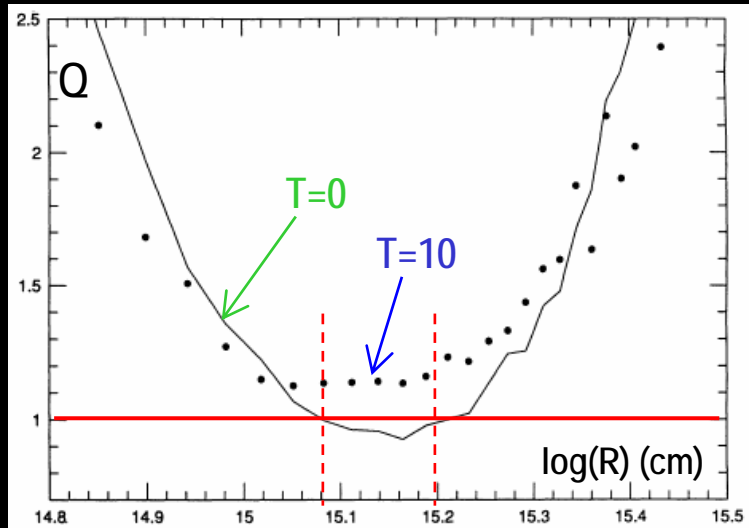
$M_{core} \sim 5 M_{\oplus}$ for Jupiter!

(Galileo data)

Solution = smaller dM_{plan}/dt and opacities?

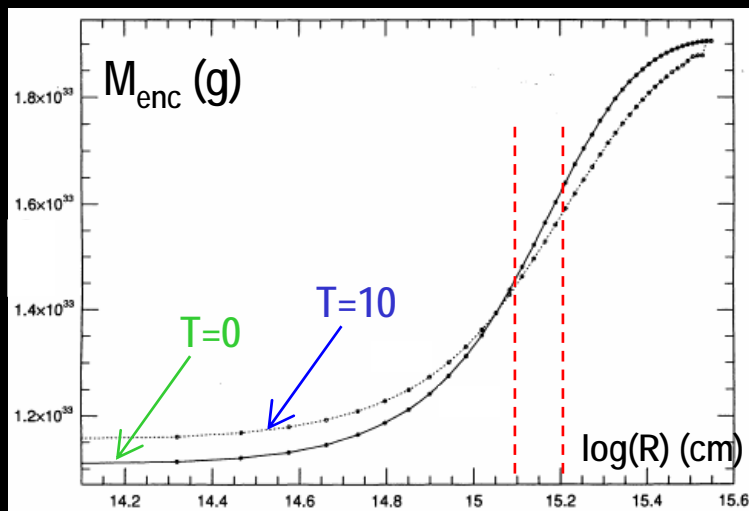
(Papaloizou & Terquem 1999)

Gravitational instabilities



Important when $Q = \frac{\kappa c}{\pi G \Sigma} \approx \frac{M_*}{M_d(R)} \frac{H}{R} \approx 1$

i.e. $M_d(R) \sim 0.1 M_*$



Instability timescale: a few Ω^{-1}

Probably help pushing a significant amount of mass onto the star in the early stages.

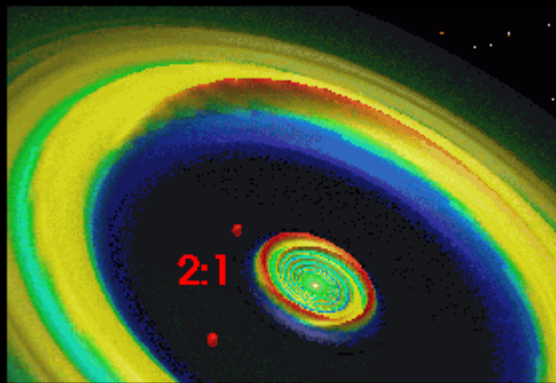
(Laughlin & Bodenheimer 1994)

Planet migration



(L. Cook)

Hot Jupiter and Neptunes:
in situ formation:
too hot
(not enough material)

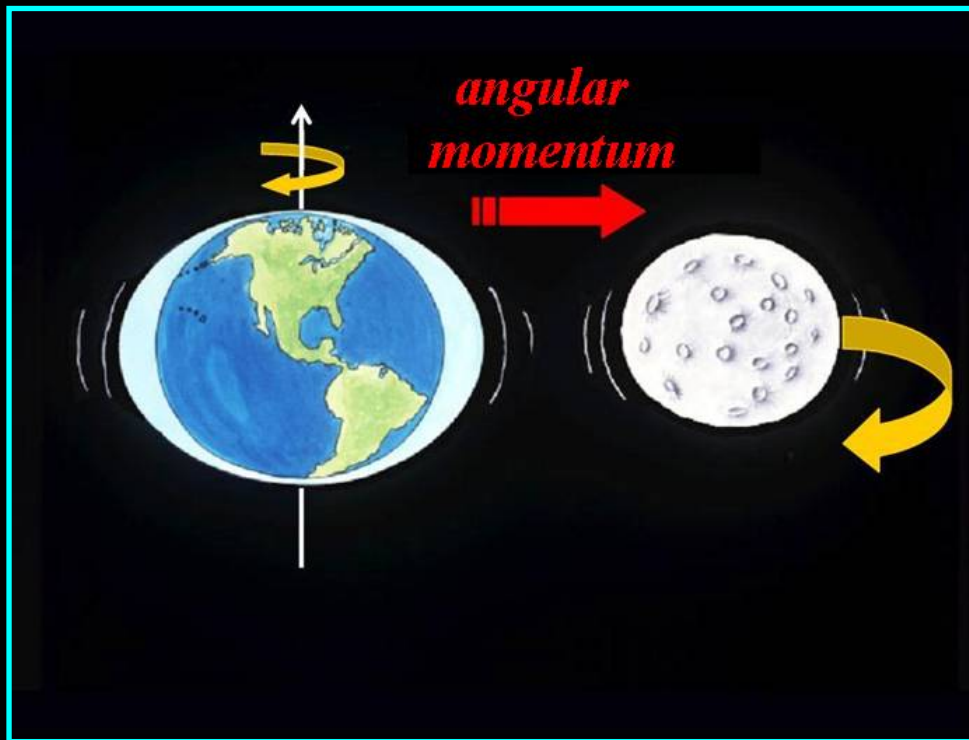


(G. Bryden)

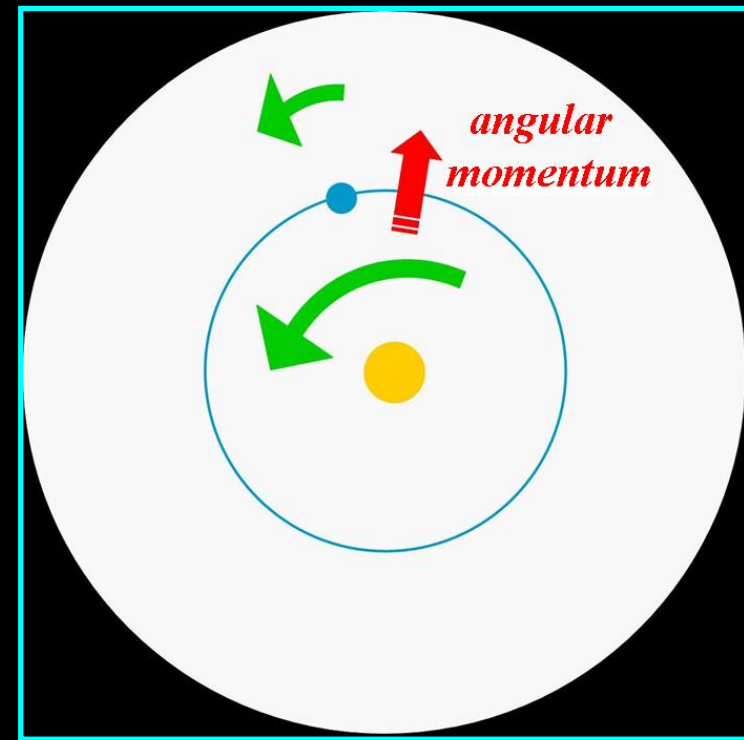
Resonant planets:
capture

Tidal torques

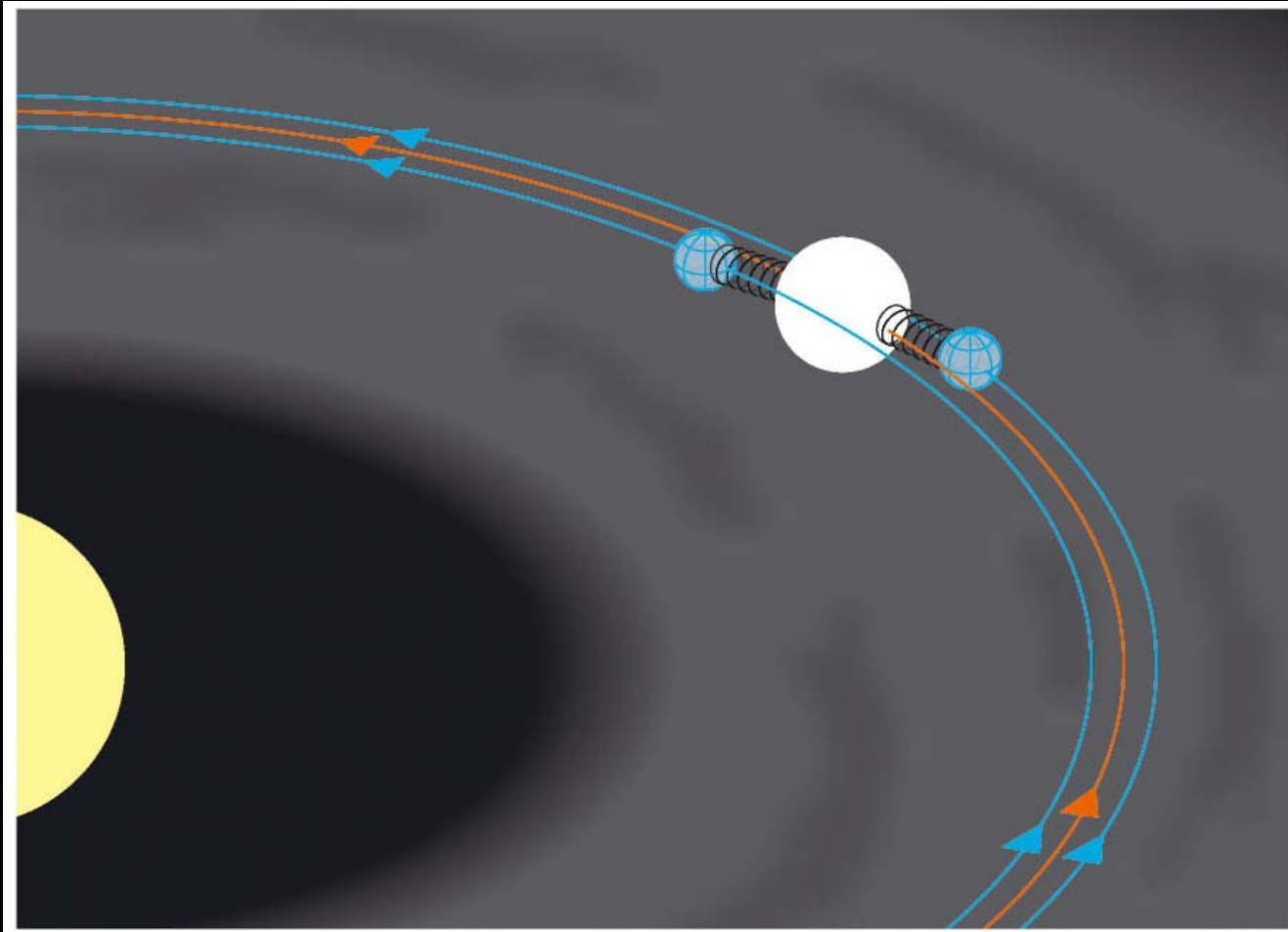
Earth-Moon system



Keplerian disk



Tidal torques

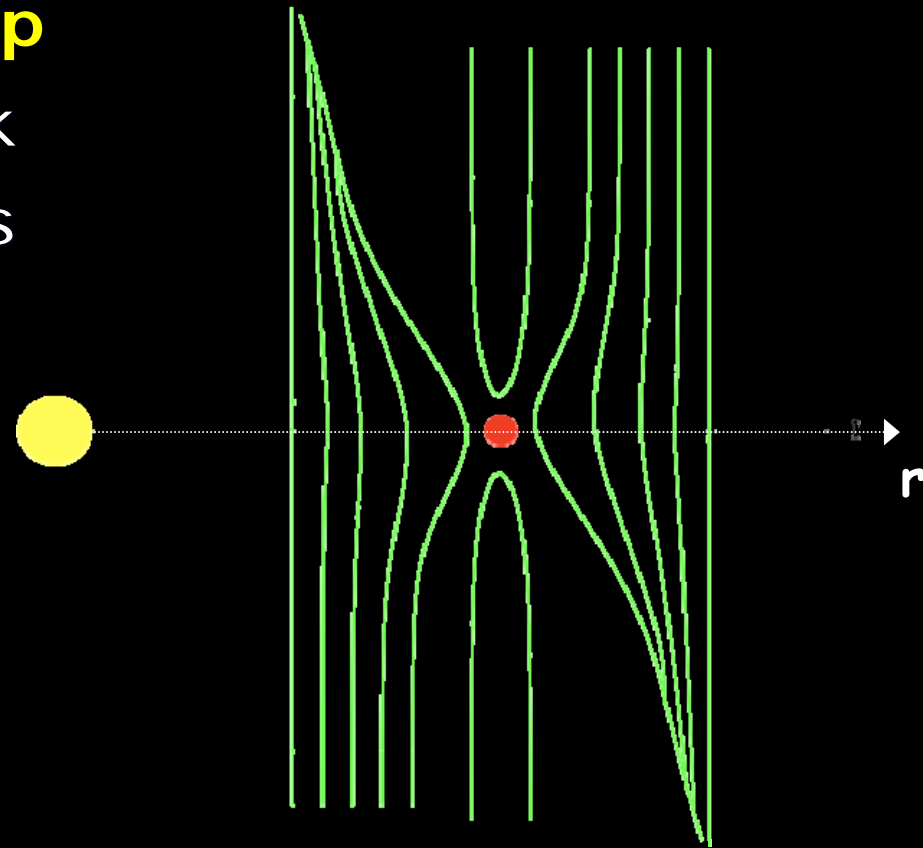


(Goldreich & Tremaine '79)

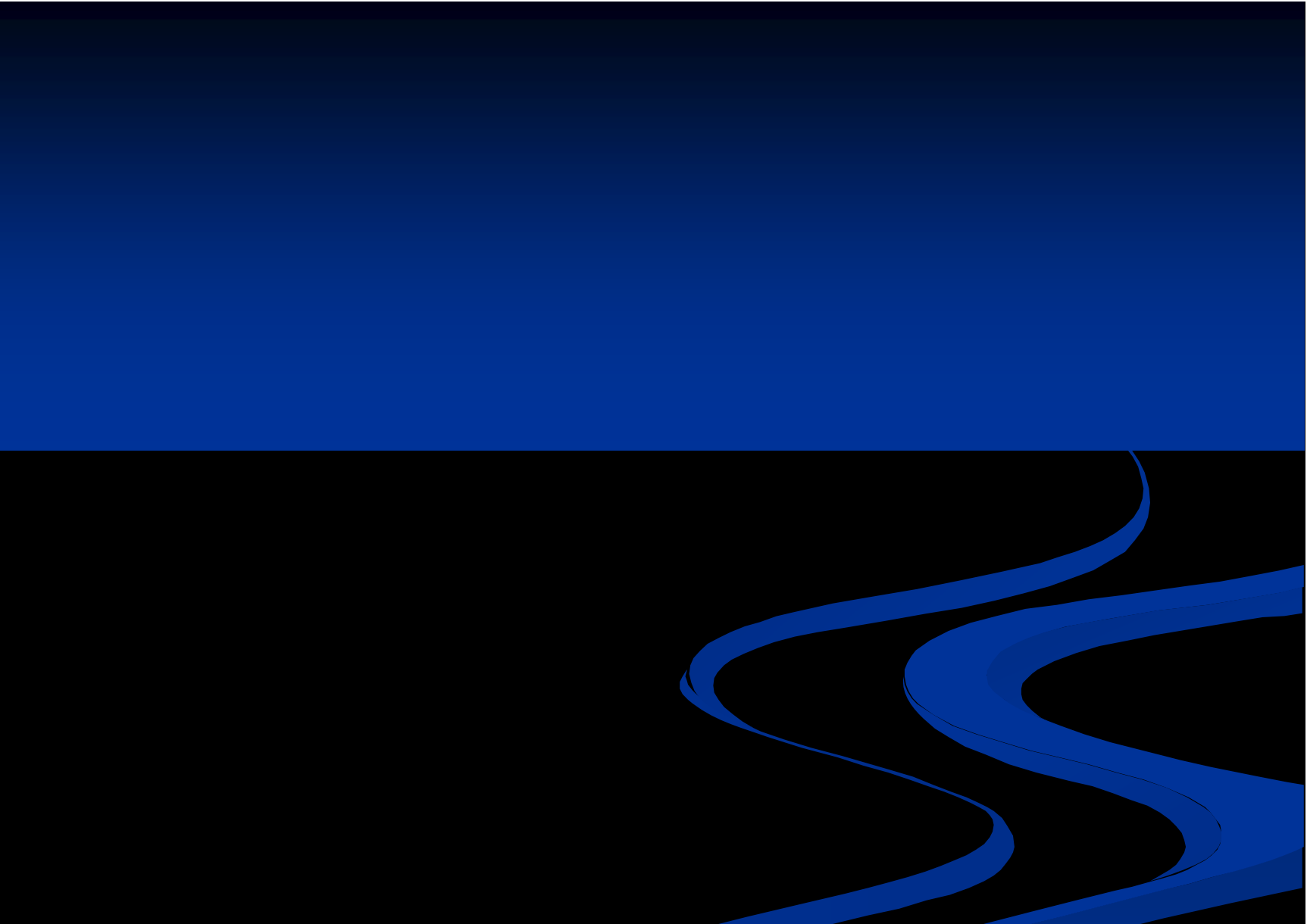


Type II migration ($M > 100 M_{\oplus}$)

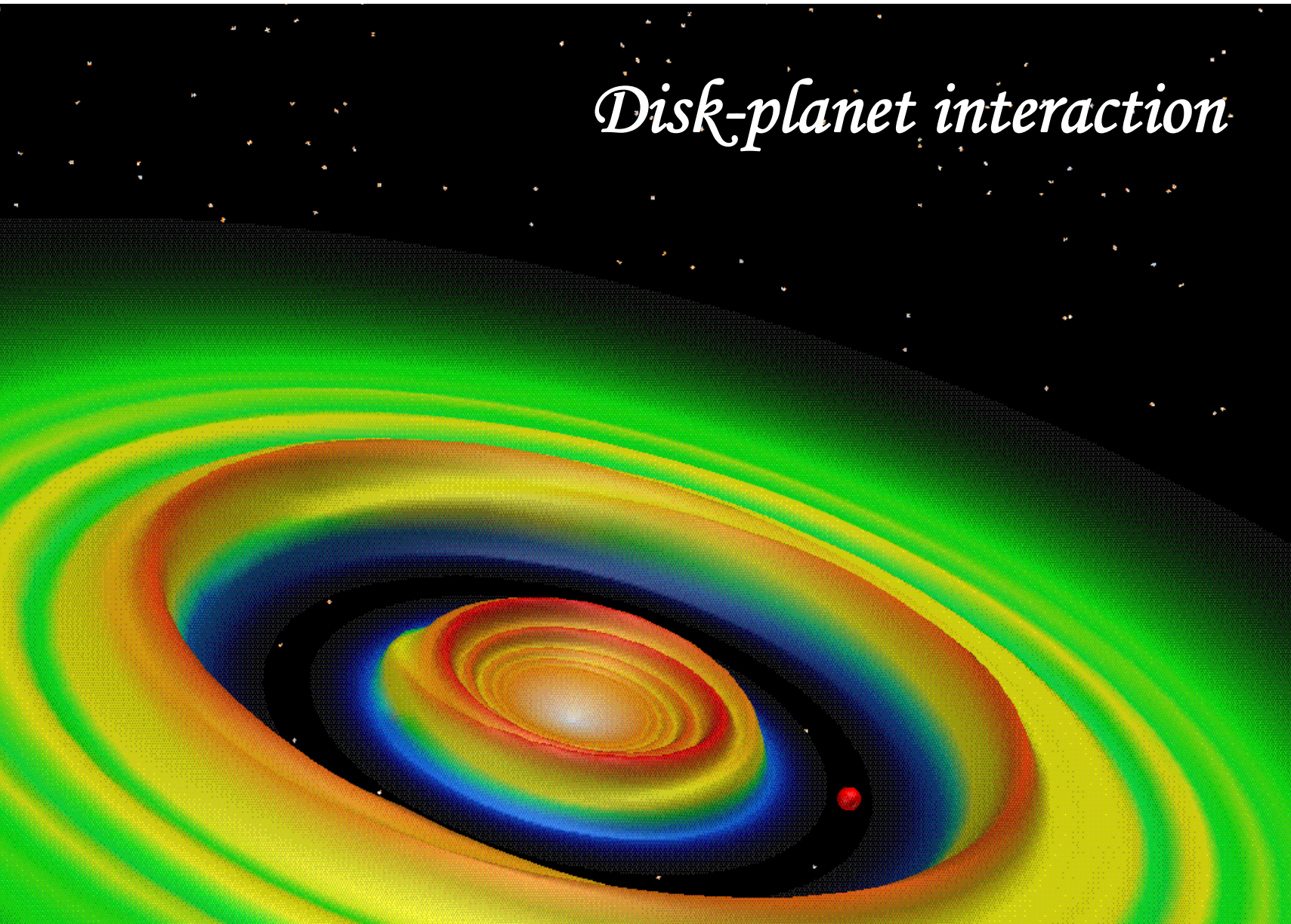
The planet opens up a **gap** and is locked into the disk viscous evolution process



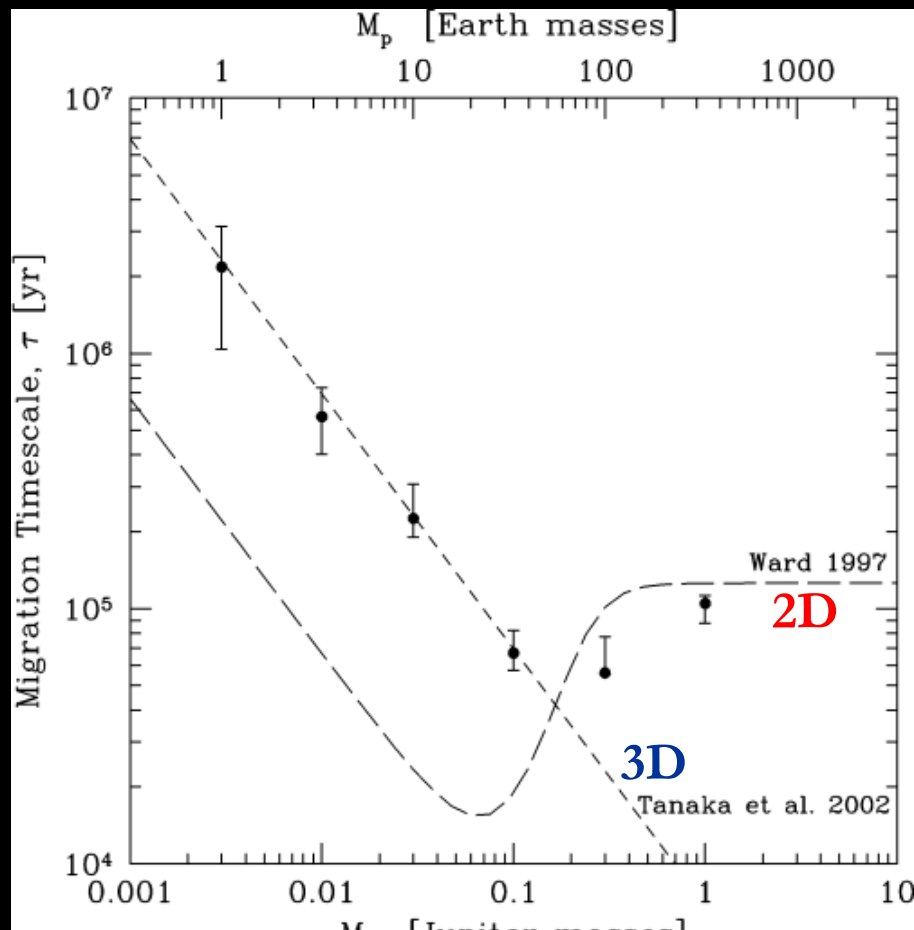
(Goldreich & Tremaine '80, Papaloizou & Lin '84)



Disk-planet interaction



Migration rate



Is migration inevitable?

Type I migration:

may not be that efficient (turbulence, magnetic field, eccentric orbits...)

Type II migration:

depends on the disk mass: efficient only in massive disks?

Conclusions

Planet formation:

Still lots of uncertainties (planetesimal formation, timescales, critical core mass)

Migration:

How much? and what about the solar system?

We do understand mechanisms, but we still lack a global view

Progress in theory and modelling is fast, but observations progress even faster!